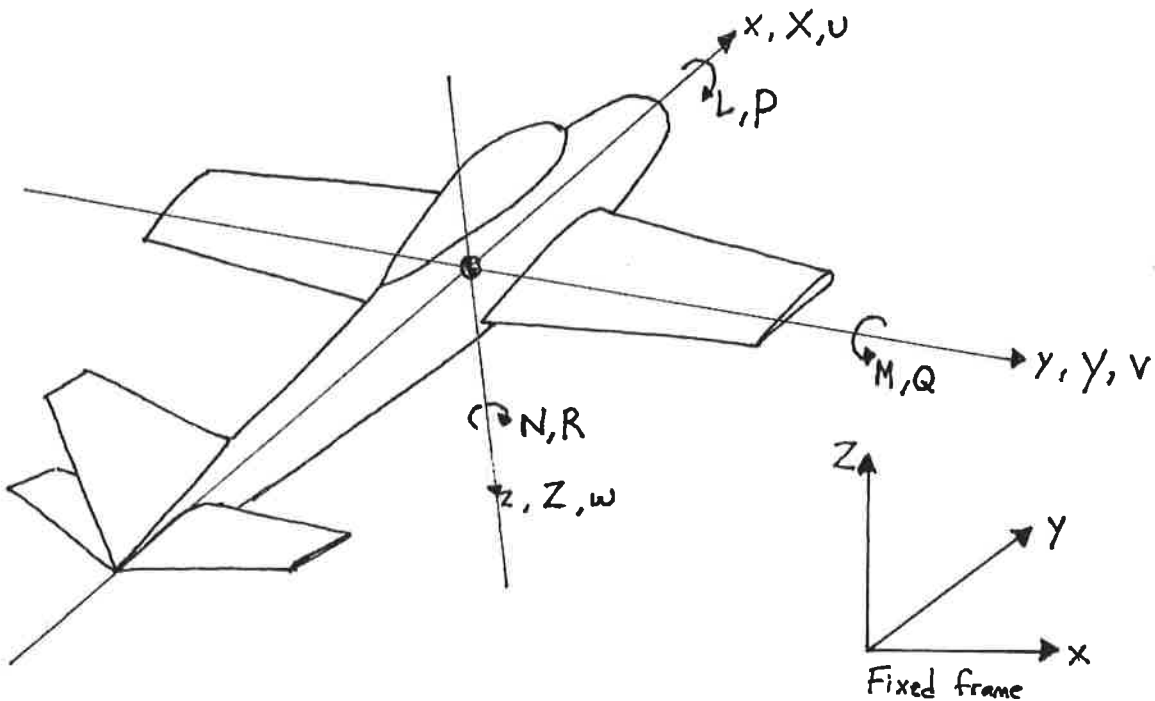


AEM 368

Flight Dynamics and Control 1

Stability and Control Experiment

Aircraft Coordinate System



- x, y, z Aircraft "stability" frame location
- X, Y, Z Forces in stability frame
- u, v, w Velocity in stability frame
- L, M, N moment in stability frame
- P, Q, R Angular velocities (roll, pitch, yaw)
- ϕ, θ, ψ Euler angles (orientation)

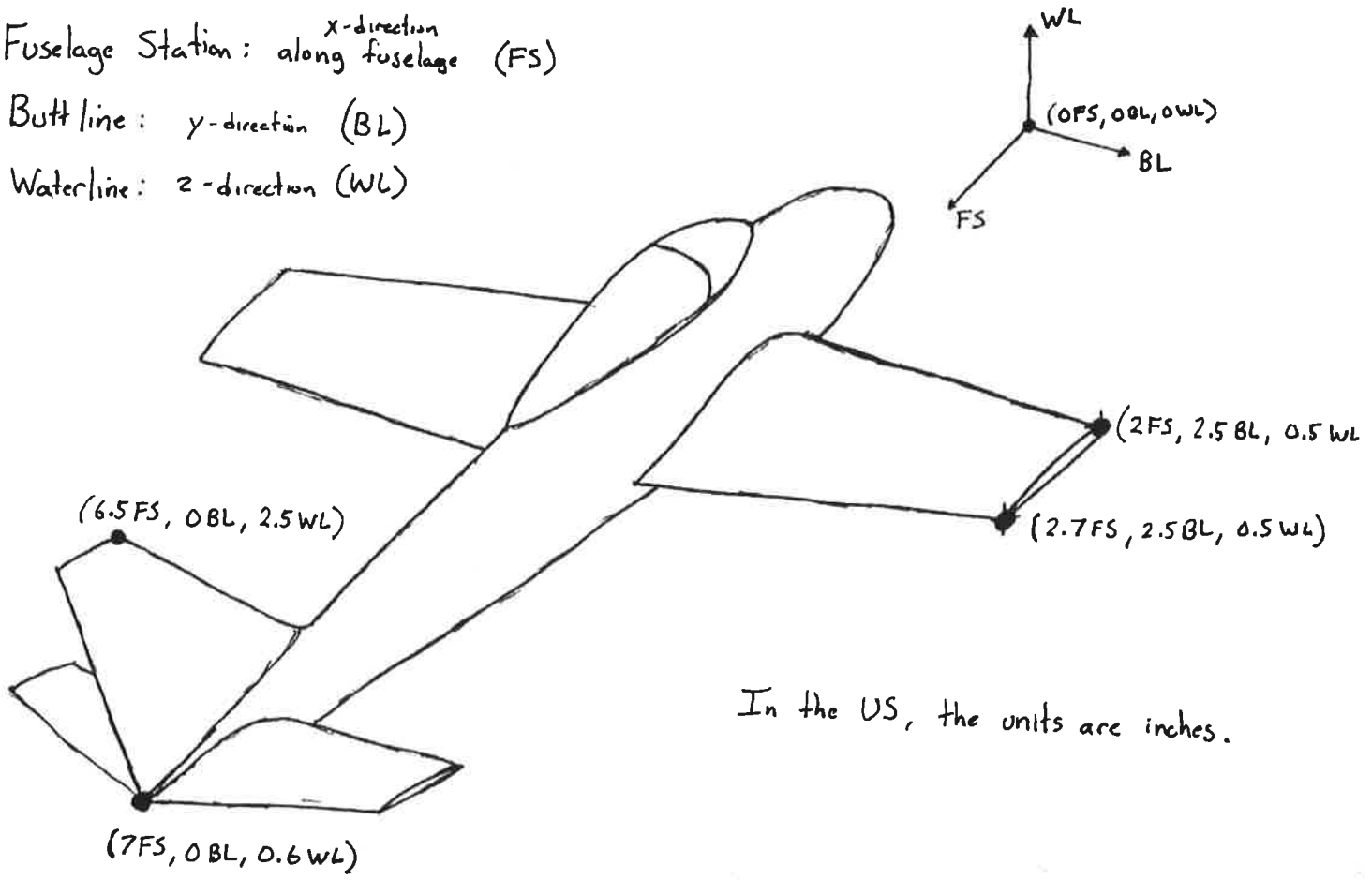
Warning:

Outside of aerodynamics, the x axis is usually down the length of the a/c.

$$\begin{aligned} X_{\text{loft}} &= -X_{\text{aero stability}} \\ Y_{\text{loft}} &= Y_{\text{aero stability}} \\ Z_{\text{loft}} &= -Z_{\text{aero stability}} \end{aligned}$$

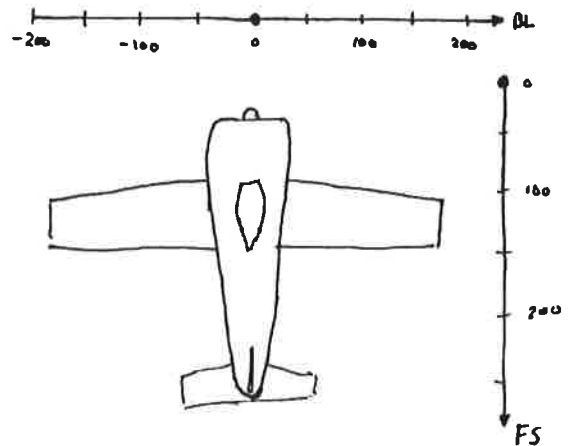
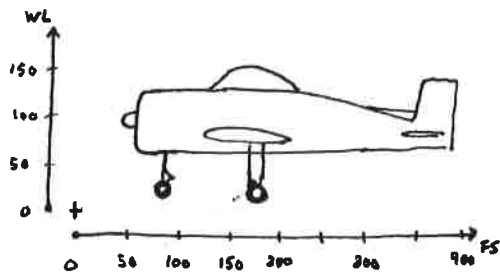
Locations on an Aircraft (Aircraft Station Coordinates)

- Fuselage Station: ^{x-direction} along fuselage (FS)
- Butt line: y-direction (BL)
- Waterline: z-direction (WL)

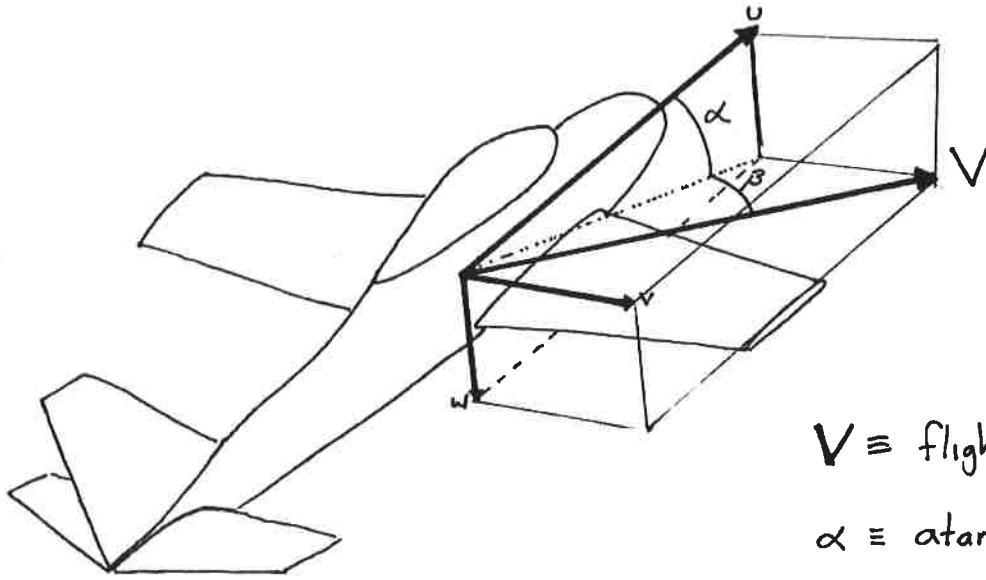


In the US, the units are inches.

- Zero FS is usually not the most forward location of the aircraft. Rather the origin is placed arbitrarily forward such that negative FS does not occur.
- Zero BL is usually along the centerline
- Zero WL is usually placed such that all values are positive



Angle of Attack and Sideslip



$V \equiv$ flight velocity vector

$\alpha \equiv \arctan\left(\frac{w}{u}\right)$ Angle of Attack

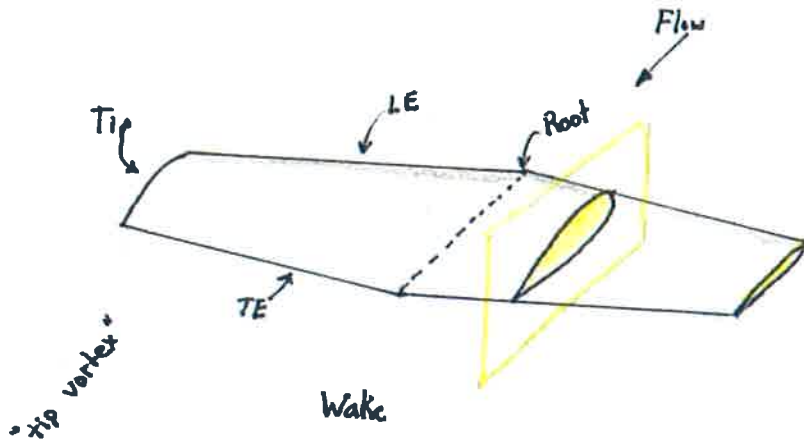
$\beta \equiv \arcsin\left(\frac{v}{V}\right)$ Sideslip

α is defined wrt the projection of V onto the body frame (i.e. U)

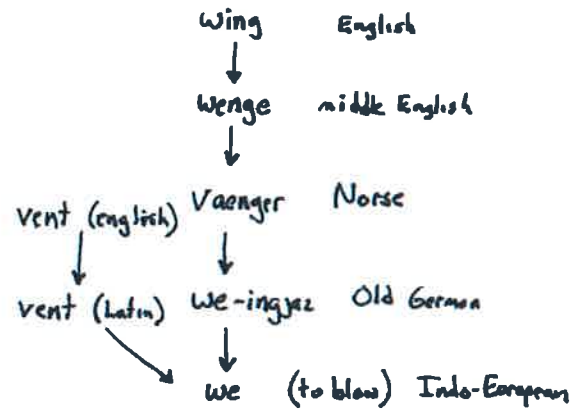
β is defined wrt the v projection and V .

Wing:

- Three dimensional closed surface
- Generates aerodynamic force
- Cross sections are airfoils

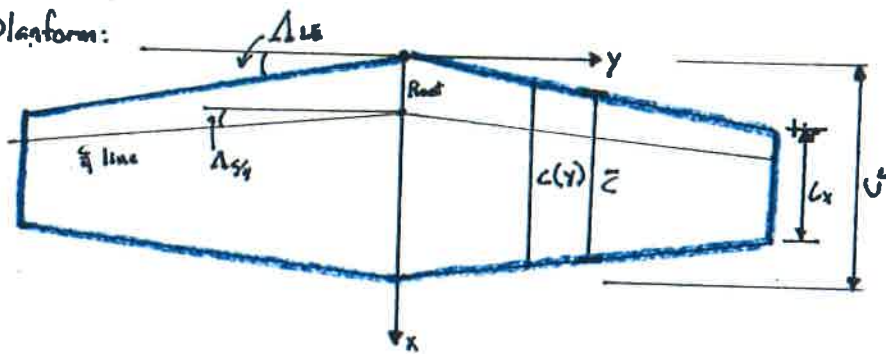


Etymology:



Interestingly, not from Indo-European word for "fly" which is "peth" from which we get "feather" and "pen".

Planform:



$S \equiv$ Wing Area $[L^2]$

$b \equiv$ Span $[L]$

$c \equiv$ chord

$AR \equiv$ Aspect Ratio $\approx \frac{b^2}{S}$

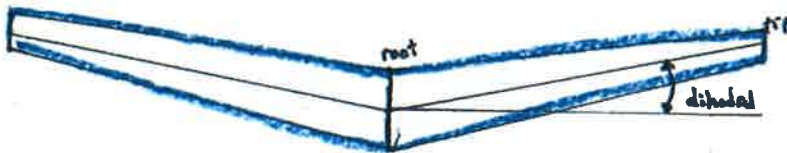
$\lambda \equiv$ taper ratio $\approx \frac{c_t}{c_r}$, $\left(\frac{tip}{root}\right)$

$MAC \equiv$ Mean Aerodynamic Chord

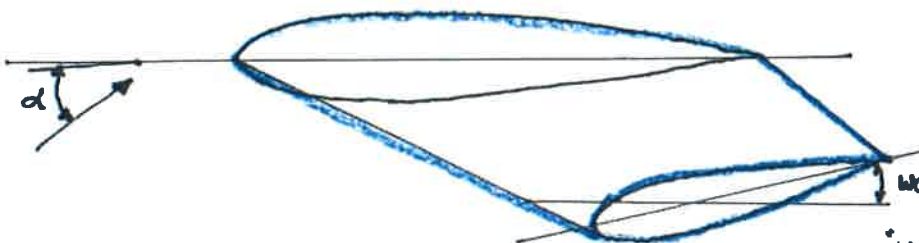
$$\frac{1}{S} \int_{-b/2}^{b/2} c(y)^2 dy$$

$\bar{c} \equiv$ Average Chord

Front view



Side view



wash out \equiv decrease in incidence angle at tip

wash in is a positive increase ... usually a bad idea

The atmosphere on Earth

Dry Air (= Atmospheric air - H₂O - contaminants (dust, pollen, ...))

	Mole Fraction	Molecular Weight $\frac{kg}{kmol} \approx \frac{lbm}{lbmol}$	
Nitrogen	78.08%	28.02	$N \equiv N$ strong triple bond!
Oxygen	20.95%	32.0	$O = O$ double bond
Argon	0.93%	39.94	Greek αργον "inactive" Ar Ar no bond
Carbon Dioxide	0.03%	44.01	$O = C = O$ linear shape double bond
Other	0.01%		
	<hr/> 100%		

Apparent Molecular Weight = $\sum a_i M_i$

$$M \approx 28.02 \cdot 0.7808 + 32.0 \cdot 0.2095 + 39.94 \cdot 0.0093 + 44.01 \cdot 0.0003$$

$$\approx 28.97 \frac{lbm}{lbmol} = 28.97 \frac{kg}{mol}$$

Gas Constant for air

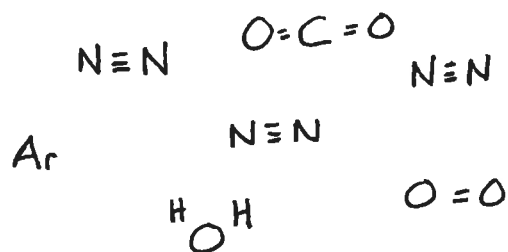
$$R = \frac{\bar{R}}{M} = \frac{1545.34 \frac{ft \cdot lbf}{R \cdot lbmol}}{28.97 \frac{lbm}{lbmol}} = \frac{1716.5 \frac{ft \cdot lbf}{lbmol \cdot R}}{28.97 \frac{lbm}{lbmol}} = 53.35 \frac{ft \cdot lbf}{lbm \cdot R}$$

Air density SSL (14.696 psi, 59°F)

$$p = \frac{pM}{\bar{R}T} = \frac{14.696 \text{ psi} \cdot 28.97 \frac{lbm}{lbmol}}{1545.34 \frac{ft \cdot lbf}{lbmol \cdot R} \cdot 519.67 \text{ R}} = 0.00237 \frac{slug}{ft^3}$$

Wet air

The addition of water vapor changes the properties of "air":



Water vapor behaves as an ideal gas, thus we can model the mixture as an IG.

$$p = \frac{P_{\text{dry}}}{R_{\text{dry}} T_{\text{dry}}} + \frac{P_{\text{vapor}}}{R_{\text{vapor}} T_{\text{vapor}}}$$

where the partial pressures add to total pressure
 $P = P_{\text{dry}} + P_{\text{vapor}}$

$$= \frac{P_{\text{dry}} M_{\text{dry}}}{R T_{\text{dry}}} + \frac{P_{\text{vapor}} M_{\text{vapor}}}{R T_{\text{vapor}}}$$

Temps are identical

$$= \frac{P_{\text{dry}} M_{\text{dry}} + P_{\text{vapor}} M_{\text{vapor}}}{RT}$$

Partial pressure of vapor

$$P_{\text{vapor}} = \phi P_{\text{sat}}$$

$\phi \equiv$ relative humidity

$$= \frac{(P - P_{\text{vapor}}) M_{\text{dry}} + P_{\text{vapor}} M_{\text{vapor}}}{RT}$$

$$= \frac{(P - \phi P_{\text{sat}}) M_{\text{dry}} + \phi P_{\text{sat}} M_{\text{vapor}}}{RT}$$

$$= \frac{P M_{\text{dry}} + \phi P_{\text{sat}} (M_{\text{vapor}} - M_{\text{dry}})}{RT}$$

$$M_{\text{dry}} = 28.97$$

$$M_{\text{vapor}} = 18.0$$

Thus, $M_{\text{vapor}} - M_{\text{dry}}$ is negative

Increasing the water vapor decreases air density

Impact of Wet Air.

	Std-Day	Temp	rh	Altitude	Density ρ	%SSL
• Standard Sea Level (SSL)	Std-Day	59°F	0% rh	0 ^{ft} MSL	$\rho = 0.00237 \frac{\text{slug}}{\text{ft}^3}$	100
• Alabama Summer (hot + humid)		90°F	90% rh	$\approx 0^{\text{ft}}\text{MSL}$	$\rho = 0.002206 \frac{\text{slug}}{\text{ft}^3}$	93%
• " " Dry		90°F	0% rh		$\rho = 0.00224 \frac{\text{slug}}{\text{ft}^3}$	94%
• Alabama Winter (Wet)		40°F	90% rh		$\rho = 0.00246 \frac{\text{slug}}{\text{ft}^3}$	104%
• " " Dry		40°F	0% rh		$\rho = 0.00246 \frac{\text{slug}}{\text{ft}^3}$	104%
• Antarctica (cold + dry)		-126 °F	0% rh		$\rho = 0.00369 \frac{\text{slug}}{\text{ft}^3}$	155%
• Denver, CO (std day)		40°F	0% rh	5000 ^{ft} MSL	$\rho \approx 0.00205 \frac{\text{slug}}{\text{ft}^3}$	86%

ASHRAE Chart Comparison



ASHRAE PSYCHROMETRIC CHART NO.1

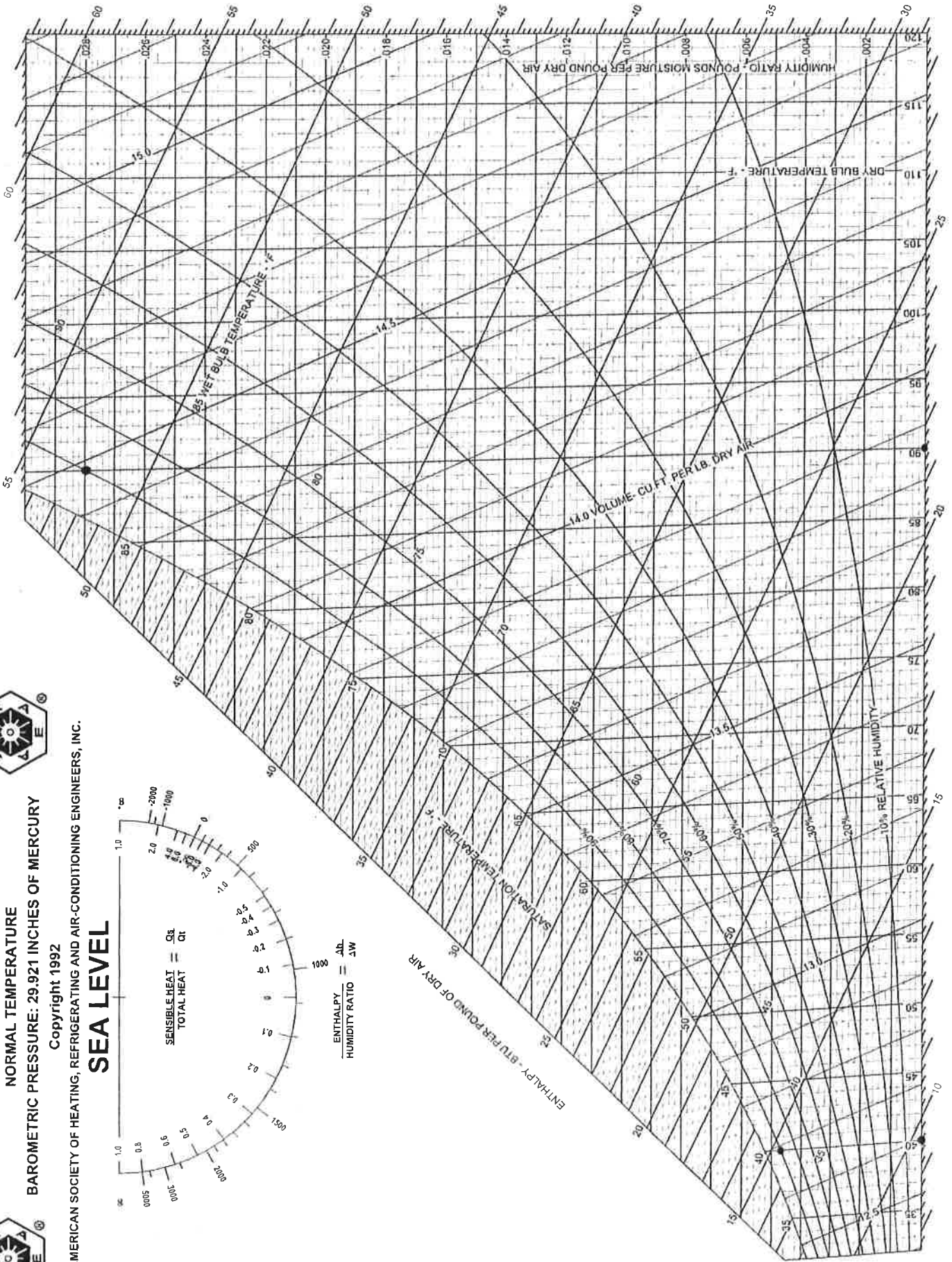
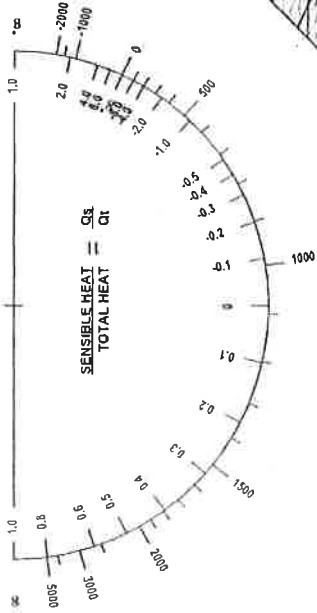
NORMAL TEMPERATURE

BAROMETRIC PRESSURE: 29.921 INCHES OF MERCURY

Copyright 1992

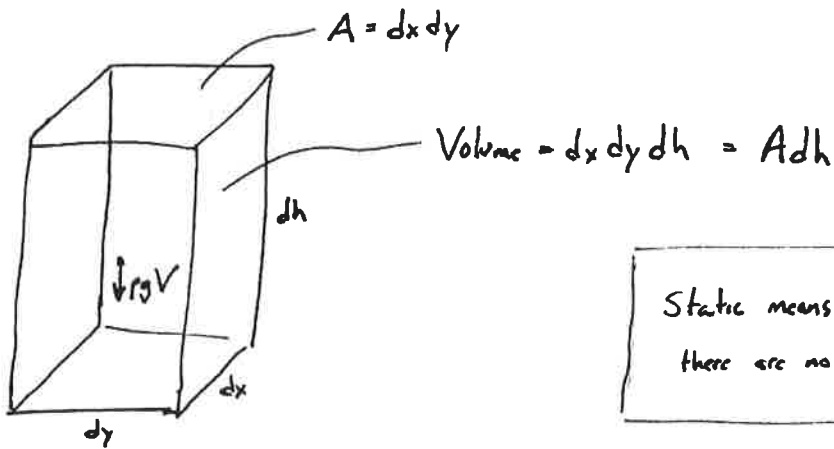
AMERICAN SOCIETY OF HEATING, REFRIGERATING AND AIR-CONDITIONING ENGINEERS, INC.

SEA LEVEL



ENTHALPY - BTU PER POUND OF DRY AIR

Static Column of Fluid



Static means no velocity, so no $\frac{du_i}{dx_j}$, thus there are no viscous forces.

Summation of forces in h direction

$$P_{\text{bottom}} A = P_{\text{top}} A + \rho g V$$

Taylor Series expansion for $P_{\text{top}} = P_{\text{bottom}} + \frac{dP}{dh} dh$

$$PA = \left(P + \frac{dP}{dh} dh \right) A + \rho g A dh$$

Reduce (divide by A , cancel PA terms)

$$\frac{dP}{dh} dh + \rho g dh = 0$$

Divide by dh

$$\frac{dP}{dh} + \rho g = 0$$

Gov Egu

$$dP = -\rho g dh$$

Atmosphere (continued)

$$dp = -\rho g_0 dh$$

Ideal Gas is $P = \rho RT$, substitute for ρ

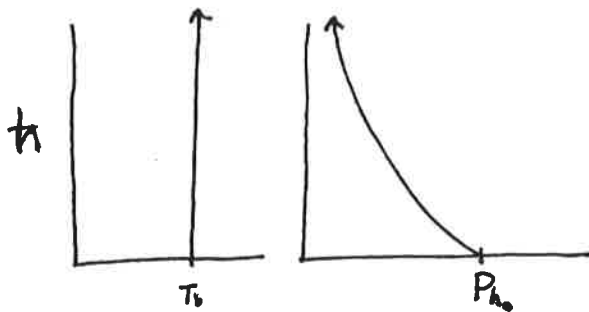
$$dp = -\frac{P}{RT} g_0 dh \Rightarrow \frac{dp}{P} = -\frac{g_0 dh}{RT}$$

$$R = 1716.5 \frac{\text{ft} \cdot \text{lb}_f}{\text{R slug}}$$

- Isothermal
Zero lapse rate ($\Delta T \neq f(h)$) ($T = T_0$)

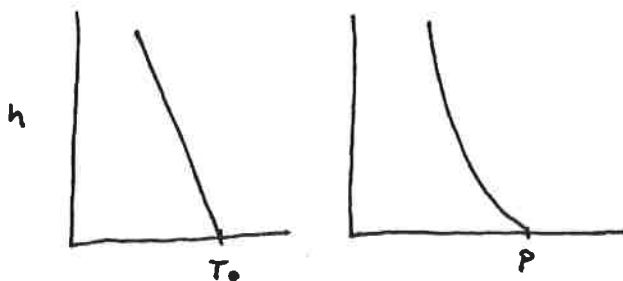
$$\frac{dp}{P} = \frac{-g_0}{RT_0} dh \quad \text{Integrate} \quad \ln \frac{P}{P_0} = \frac{-g_0}{RT_0} h \Big|_{h_0}^{h_1}$$

$$\underbrace{\ln P_1 - \ln P_0}_{\ln \frac{P_1}{P_0}} = \frac{-g_0}{RT_0} (h_1 - h_0) \Rightarrow P_1 = P_0 e^{-\frac{g_0 (h_1 - h_0)}{RT_0}}$$



- Linear lapse rate ($T = T_0 - \lambda (h_1 - h_0)$)

$$\frac{dp}{P} = \frac{-g_0}{R(T_0 - \lambda (h - h_0))} dh \quad \text{Integrate (slightly involved)} \quad \frac{P}{P_0} = \left(\frac{T}{T_0}\right)^{\frac{-g_0}{R\lambda}}$$



Lesson 2

Non standard Atmosphere

MIL-STD-210A

Hydrostatic Equation (review)

$$dp = -\rho g_0 dh$$

p is pressure
 h is geopotential altitude
 g_0 is gravity
 ρ is density

With the ideal gas equation of state

$$p = \rho R T$$

So,

$$dp = -\frac{\rho}{RT} g_0 dh$$

But $T = T(h)$ (a function of height)

We could solve analytically (as in lesson 1) or numerically.

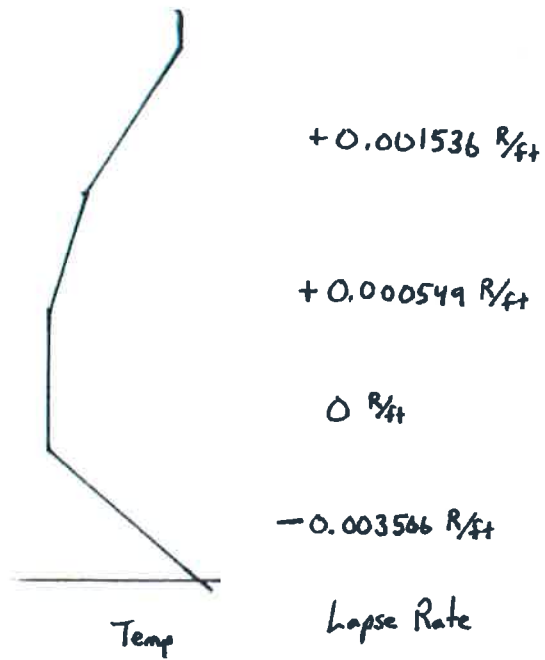
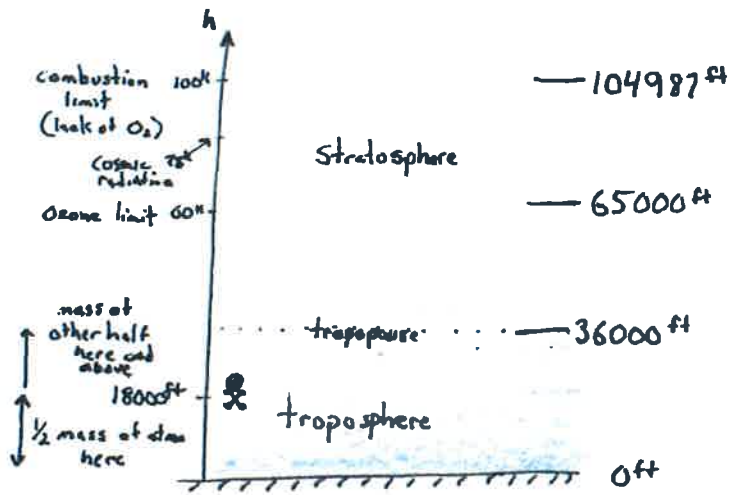
$$dp = -\frac{\rho}{RT} g_0 dh \Rightarrow p(h_1) = p(h_0) - \underbrace{\frac{p(h_1)p(h_0)}{R(T(h_1)-T(h_0))}}_{\text{one possible integration method}} g_0 (h_1 - h_0)$$

Units:

$$dp = -\frac{\rho}{RT} g_0 dh$$

$$\frac{\frac{\text{lb}_f}{\text{in}^2}}{\text{dp}} = \frac{\frac{\text{lb}_f}{\text{in}^2}}{p} \cdot \frac{R \text{ slug}}{1716.5 \frac{\text{ft} \cdot \text{lb}_f}{\text{lb}_f}} \cdot \frac{1}{R} \cdot \frac{32.174 \text{ ft}}{g_0} \cdot \frac{\text{ft}}{dh} \cdot \frac{\text{lb}_f \text{ slug}}{\text{slug} \cdot \text{ft}} \quad \checkmark$$

Standard Atmosphere — ≈ 160000 ft



We are lucky to live at the bottom of an ocean of protective air.

- Death zone for humans begins around 20k ft or so...
- Half the mass of air is below 18k ft
- Ozone above 60k ft prevents use of outside air for humans.
- Cosmic radiation becomes significant around 75k ft.
- Normal jet combustion fails around 100k ft. (Not enough O_2)
- The positive lapse rate around 65k ft makes the atmosphere stable. Convection is minimal.

Non-dimensional P, T, ρ ratios

$$\delta = \frac{P}{P_{SSL}} \quad \theta = \frac{T}{T_{SSL}} \quad \sigma = \frac{\rho}{\rho_{SSL}} = \frac{\delta}{\theta}$$

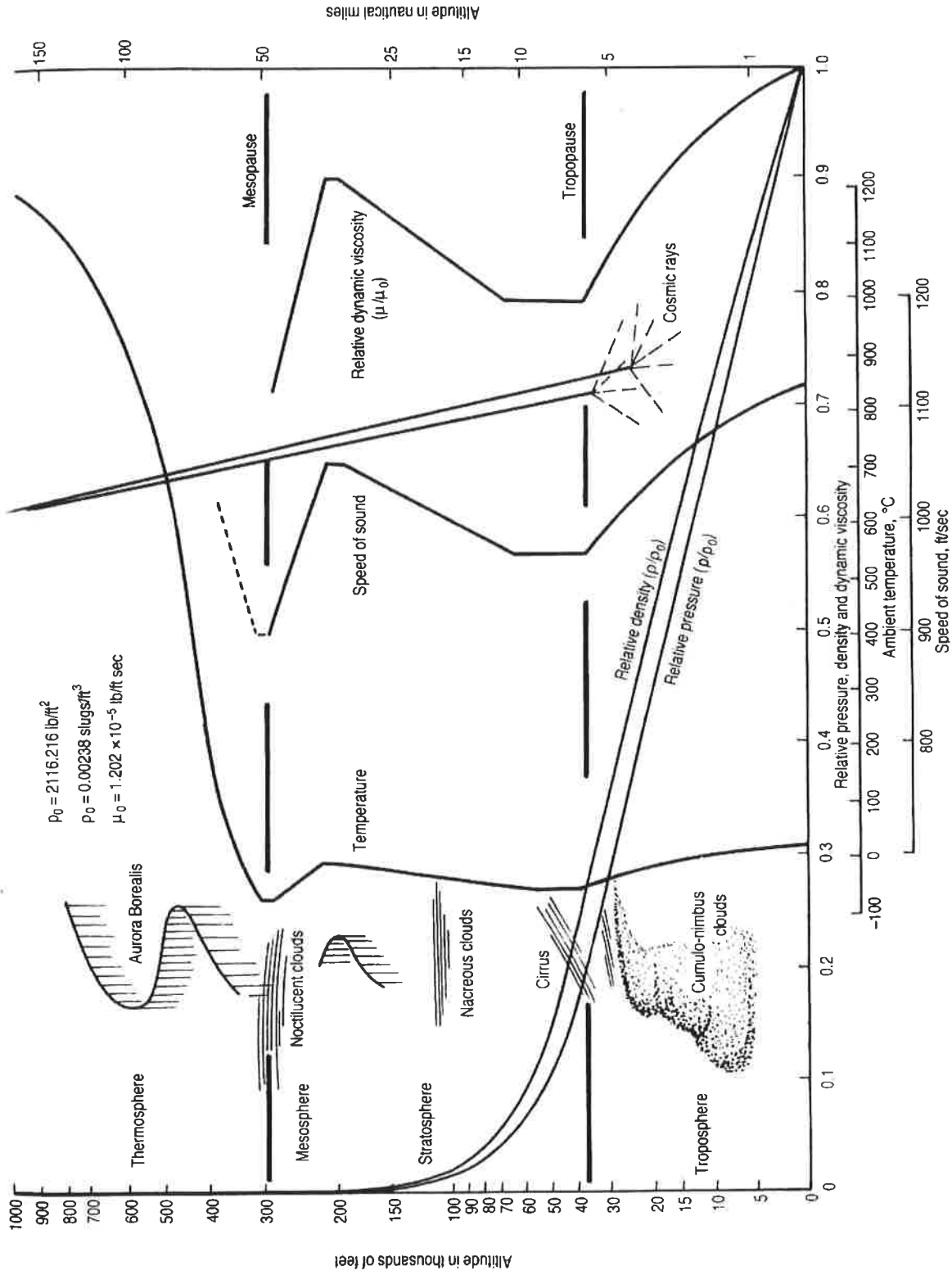
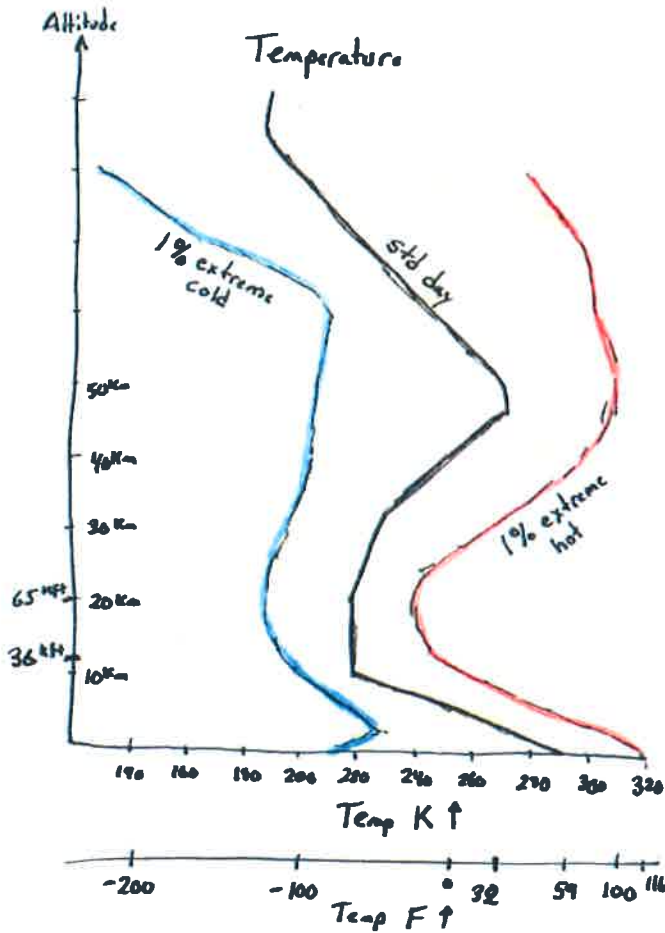


Fig. 1.2 General characteristics of the atmosphere (based upon ICAO and US Standard Atmosphere 1962).

Source: *The Anatomy of the Airplane*
 Stanton

The atmosphere is not, has not, and never will be standard.



See:

U.S. Standard Atmosphere 1976

(NASA-TM-X-74335)

for details (241 pages!)

Rankine:

$$R = F + 459.67$$

Kelvin:

$$K = C + 273.1$$

Non Std Atmosphere Models. (MIL-STD-210A)

Std h [ft]	λ [R/ft]
0	-0.003366
36089	0
65617	+0.000549
104987	+0.001536
154199	0

$T_0 = 59^\circ\text{F}$

Hot h [ft]	λ [R/ft]
0	-0.003840
39370	+0.000439
67257	+0.000768

$T_0 = 103.28^\circ\text{F}$

Tropic h [ft]	λ [R/ft]
0	-0.003840
52493	+0.002085
68898	+0.001361

$T_0 = 90.086^\circ\text{F}$

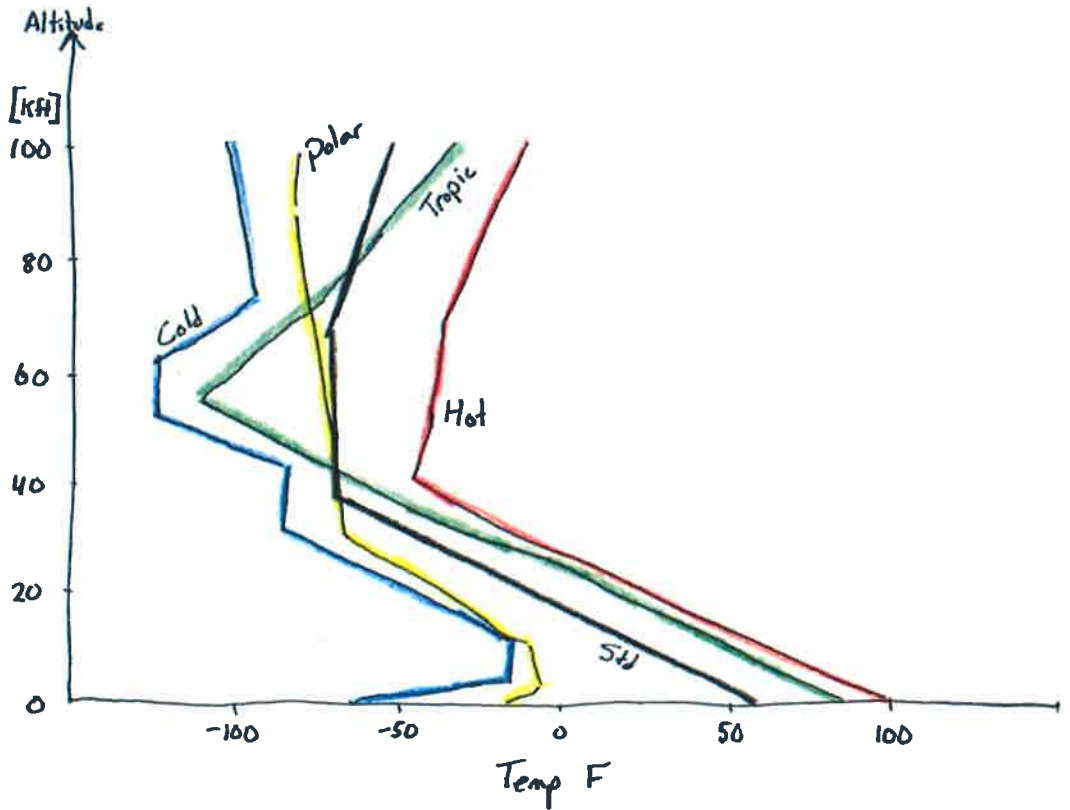
Cold	λ [R/ft]
0	+0.013716
3281	0
9843	-0.003292
31168	0
42651	+0.004872
50853	0
60696	+0.002524
73819	-0.000425

$T_0 = -59^\circ\text{F}$

Polar (Warning: reverse engineered from data!)	λ [R/ft]
0	0.003
3281	-0.00055
9843	-0.0028
31168	-0.0003
88000	0

$T_0 = -15.67^\circ\text{F}$

ISA + X
Add X to temp at
std day



Examples:

What is the temperature at 10000 ft on a standard day?

$$T(10^{kt}) = T(SSL) + \lambda \cdot h$$
$$= 59^{\circ}F - 0.003566 \frac{R}{ft} \cdot 10000 ft$$

$$T = 23.3^{\circ}F$$

$$\theta = \frac{T}{T_{SSL}} = \frac{23.3^{\circ}F + 459.67}{59 + 459.67} = 0.93 = \theta$$

What is the temperature at 10000 ft on an ISA + 20 day?

$$T_{10000ft std} = 23.3^{\circ}F$$

$$T_{10^{kt} ISA+20} = 43.3^{\circ}F$$

Just add

What is the pressure at 10000 ft on a std day?

Linear lapse rate $\lambda = -0.003566 \frac{R}{ft}$

$$\frac{P}{P_0} = \left(\frac{I}{T_0}\right)^{\frac{-g_0}{R\lambda}}$$

$$= (0.93)^{-\frac{32.174 \frac{ft}{s^2}}{1716.5 \frac{ft}{lb} \cdot \frac{slug}{ft}} \cdot \frac{ft}{-0.003566 \frac{R}{ft}} \cdot \frac{lb \cdot s^2}{slug \cdot ft}}$$

$$\rho = 0.6828$$

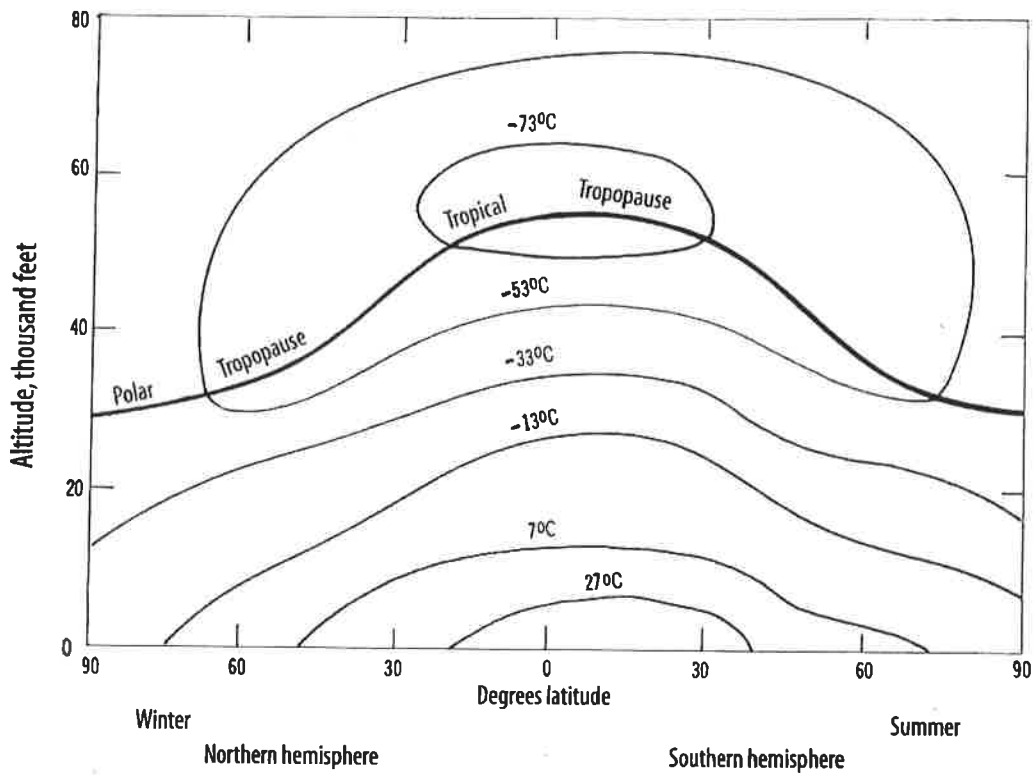


Fig. 1.1 Typical variation in atmospheric temperature along a meridian of longitude, summer in the Southern Hemisphere.

Source: TA of the Airplane, Stinson

Fact: Cold temperature testing of a/c is ^{often} done at high altitudes in the tropics!

Why? We need to know some meteorology.

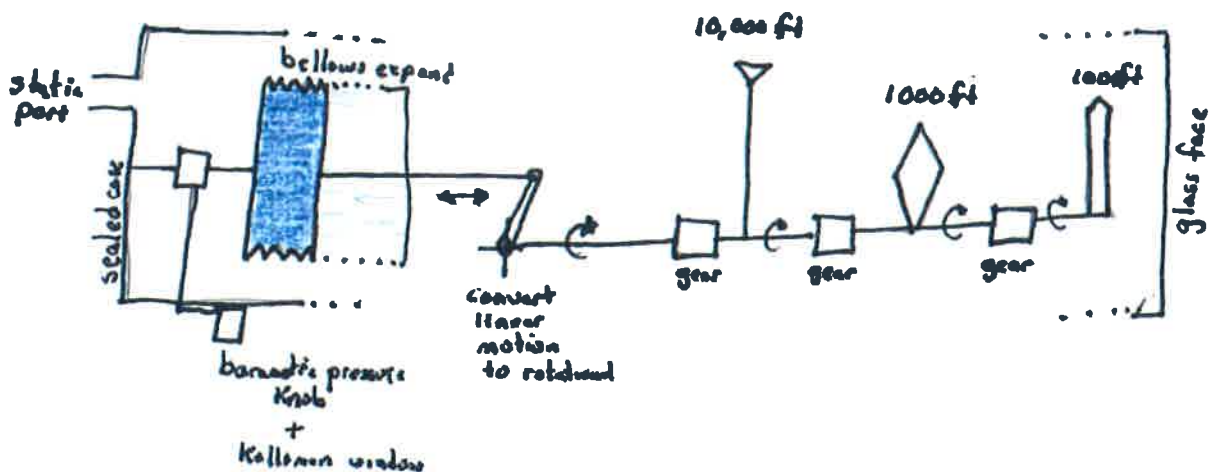
Altimeter (pressure type)

A sensor used to measure altitude. Uses local static pressure



Reading 13 700 ft

Inside (cartoon, not to scale)



Calibrated to SSL and Standard Atmosphere.

- Given a standard day, the altimeter reads a geometric altitude
- Given a non standard day, the altimeter reads pressure altitude (if 29.92 setting)
- The knob corrects for local pressure differences such that the altimeter reads the airport's elevation. Obviously, a non standard pressure profile combined with airport approach and departures makes knowing the correct Kollsman window settings critical. "Altimeter setting, two nine eight two"

Flight Level (FL):

When set to 29.92 inHg, the altimeter's reading is a convenient reference for high altitude aircraft. Read \times hundreds of feet. FL350 = 35000 ft

Q: why not low altitude?

Q: Why is a reference needed?

RVSM (Reduced Vertical Separation Minima)

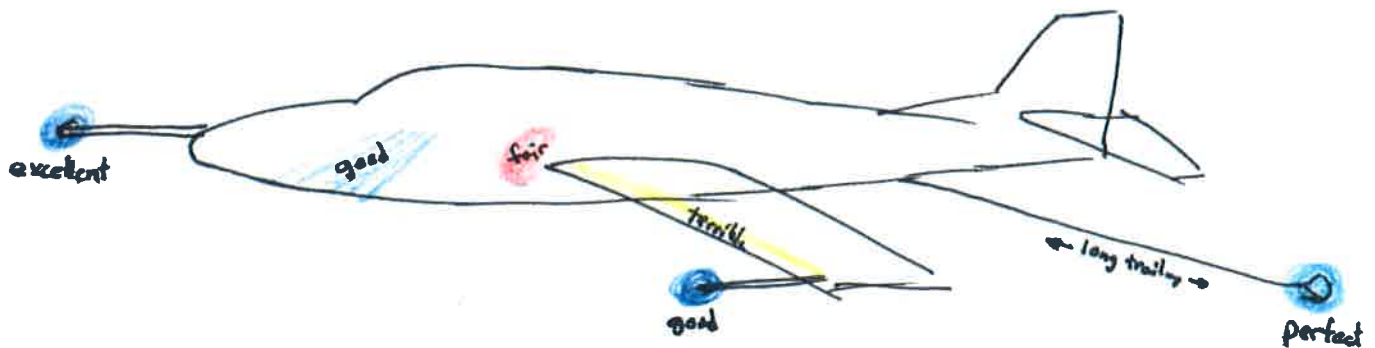
To increase the number of aircraft operating (safely) between FL 290 and FL 410, the vertical spacing between aircraft was reduced from 2000 ft to 1000 ft.

Aircraft operating in that region require special certification. In particular, the static pressure source becomes critical.

Easy Right?!? No, placing the static port(s) can be a challenge.

Given that an aircraft creates a pressure distribution to generate lift and that a fuselage impacts the pressure distribution, the static sources must be placed in a region such that minimum deviation occurs during cruise and App/Dep.
low α high α

So, where?! Usually...



Lesson 3
Aerodynamics

Lagrangian and Eulerian Frames

$$\underbrace{\frac{D(\cdot)}{Dt}}_{\substack{\text{particle} \\ \text{frame} \\ \text{"Lagrangian"}}} = \underbrace{\frac{\partial(\cdot)}{\partial t} + \mathbf{v} \cdot \nabla(\cdot)}_{\substack{\text{Eulerian frame} \\ \text{fixed location}}} = \frac{\partial(\cdot)}{\partial t} + \frac{dU_i}{dx_j} \frac{dx_j}{dt}$$

Example:

If the lapse rate is -5° per 1000ft and an aircraft is climbing at $1500\text{ft}/\text{min}$, what is the rate of change in temperature on the aircraft?

A: We want the temperature along the aircraft moving at $1500\text{ft}/\text{min}$. This is a Lagrangian frame.

$$\frac{DT}{Dt} = \frac{\partial T}{\partial t} + \frac{dT}{dh} \cdot \frac{dh}{dt} = 1500 \frac{\text{ft}}{\text{min}} \cdot \frac{dT}{dh} = 1500 \frac{\text{ft}}{\text{min}} \cdot \frac{-5^\circ}{1000\text{ft}} = -7.5^\circ/\text{min}$$

$\underbrace{\frac{\partial T}{\partial t}}_{\substack{\text{no change} \\ \text{in temp} \\ \text{at a const} \\ \text{height w/ time}}}$

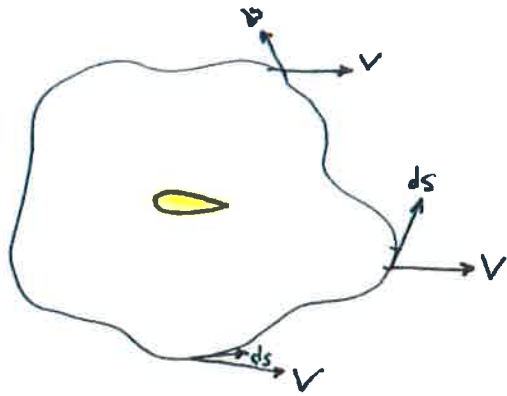
Q: If the atmosphere is now cooling at $1^\circ/\text{min}$, what is the rate of change in temperature on the aircraft?

$$\frac{DT}{Dt} = \frac{\partial T}{\partial t} + \frac{dT}{dh} \cdot \frac{dh}{dt} = -1^\circ/\text{min} + \frac{-5^\circ}{1000\text{ft}} \cdot \frac{1500\text{ft}}{\text{min}} = -8.5^\circ/\text{min}$$

Use the frame that simplifies the problem



Circulation and Lift



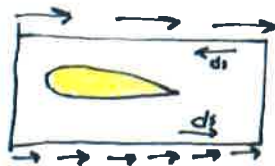
$$\text{Circulation} \equiv \Gamma = - \oint V \cdot ds$$

proportional to the velocity component tangent to a closed curve

Lift is $L = \rho V \Gamma$ ← circulation:
 ↳ density ↳ freestream velocity

No circulation no lift.

So on average, an airfoil generating lift has a net negative $V \cdot ds$. One way this can happen is if the flow above the airfoil is moving faster than the flow below

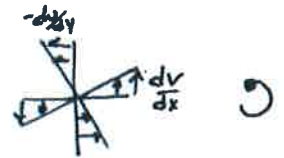


$$\left. \begin{array}{l} V \cdot ds \text{ is negative and larger} \\ V \cdot ds \text{ is positive and smaller} \end{array} \right\} L > 0$$

Vorticity:

The local velocity derivatives represent the pointwise vorticity

$$\omega = \nabla \times V \quad \text{in 2D rect' coord} = \frac{dv}{dy} - \frac{dv}{dx}$$



you can measure vorticity with an object in a flow by watching the rotation.

Vorticity and Circulation are related by

$$\Gamma = - \oint V \cdot ds = - \iint \omega \cdot \hat{n} dA$$

Since Lift requires Circulation .
and

Circulation is an integrated form of vorticity

Units:

$$\frac{\left[\frac{ft}{s} \right]}{ft} = \frac{1}{s}$$

Lift requires vorticity ... somewhere

Lift

2π corrected for aspect ratio

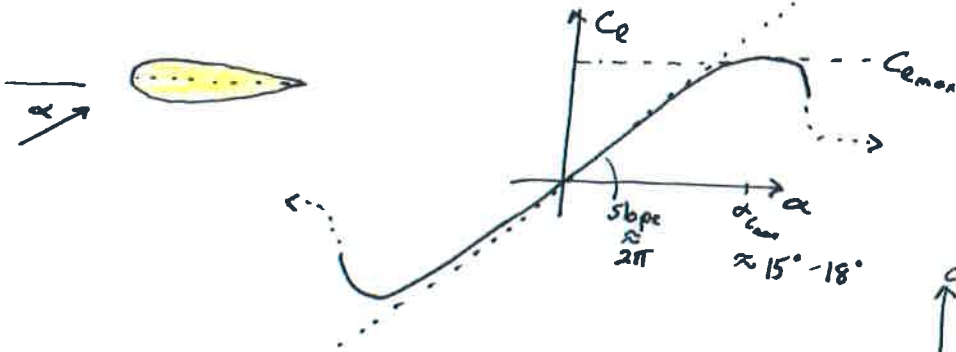
$$\text{Lift} = C_L \rho S$$

$\underbrace{\quad}_{\text{lift coeff.}}$
 $\underbrace{\quad}_{\text{dyn. press.}}$
 $\underbrace{\quad}_{\text{area}}$

2-Dimensional airfoil

$$C_{L\alpha} \approx 2\pi$$

$$C_{L\alpha} = \frac{dC_L}{d\alpha} \text{ in radians}$$



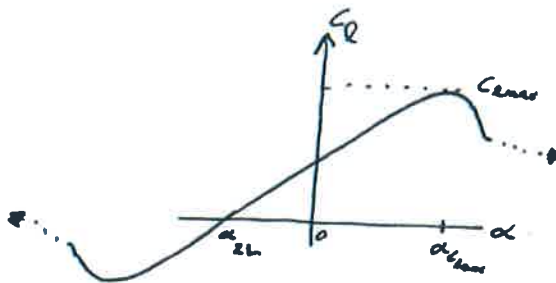
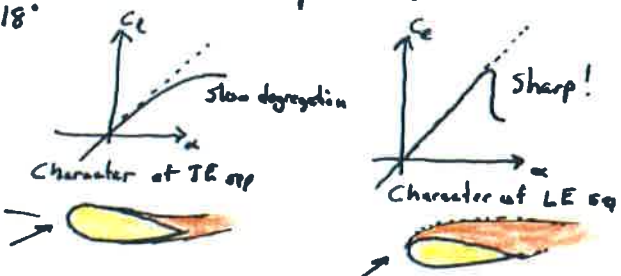
Lower C_L deviation as α increases is due to flow separation.

Camber



Mean chord line is through the forward and LE and rear TE pts.

Adding camber affects the C_L vs α curve such that zero lift occurs at α_{zL} .

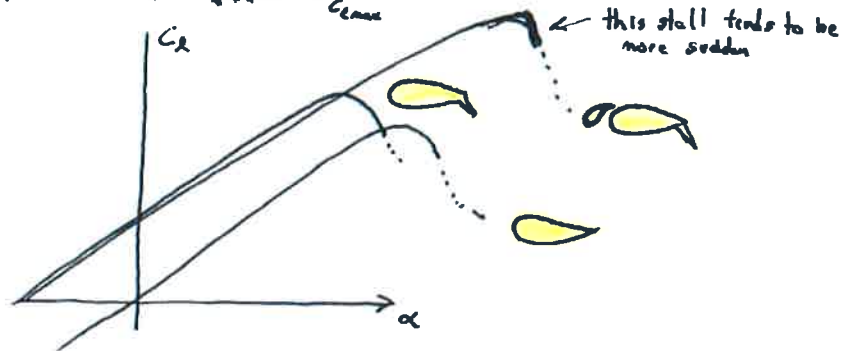


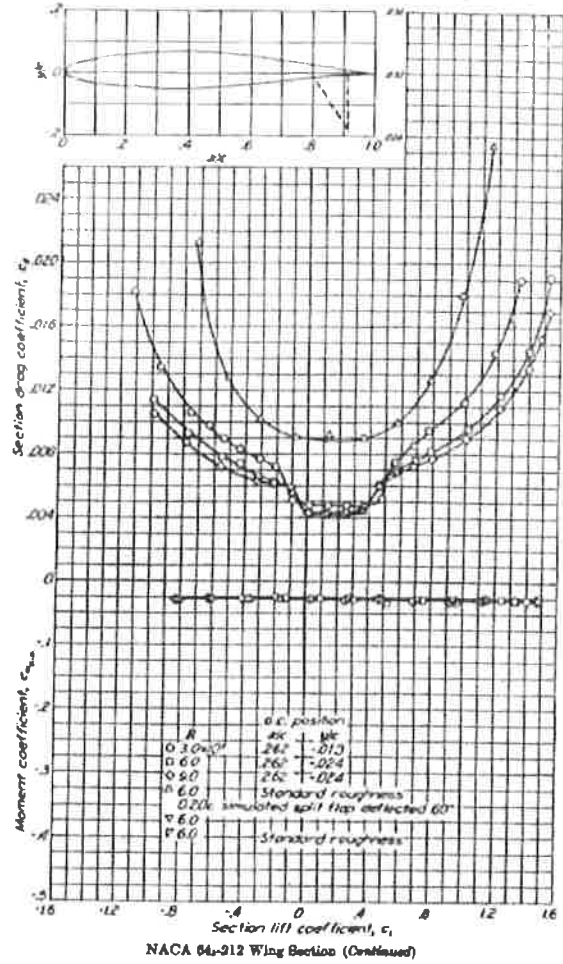
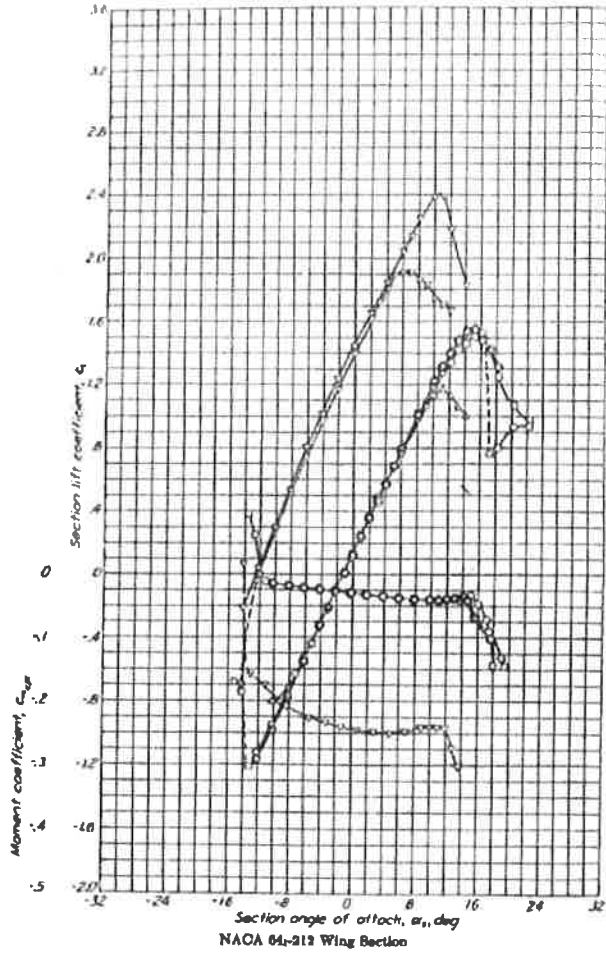
Knowing $C_{L\alpha} \approx 2\pi$ and α_{zL} , what is a good estimate for $C_L(\alpha=0)$?

$$C_L(\alpha=0) \approx 2\pi \cdot (-\alpha_{zL})$$

Slats and Flaps

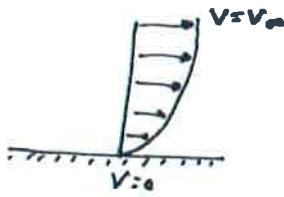
Flaps shift the C_L vs α curve upward. Slats extend α_{cmax} .





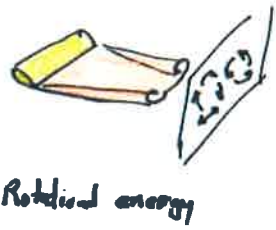
Drag

Surface Friction



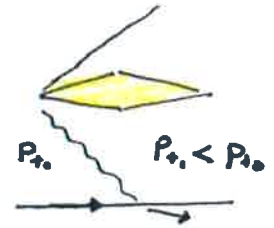
Induced Drag

+



+

Wave Drag



In general, drag is complicated to estimate for arbitrary configurations

See:

Fluid-Dynamic Drag, Hoerner

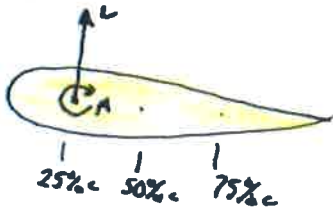
Integrated Forces and Moments on 2D airfoils.

$$L = \rho C_L S$$

$$D = \rho C_D S$$

$$M = \rho C_M S \bar{c}$$

Subsonic Airfoils



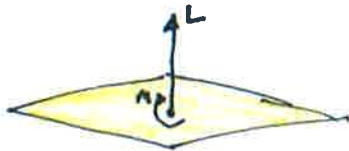
Lift and moment act at the aerodynamic center.

25% chord

$$A.C. \equiv \frac{dC_m}{dC_L} = 0$$

$$\text{Thus } C_m(\alpha) = C_{m_0}$$

Supersonic Airfoil



A.C. at 50% chord

See AEM 413

Transonic

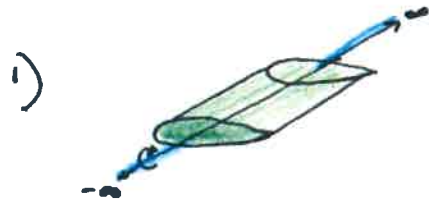
Rapid and non-monotonic shift in a.c.

Strongly depends on airfoil shape

3D Wings

From physics, an inviscid vortex must start and end on a solid surface or form a closed loop.

Possibilities:



Vortex extends off wingtips forever.

$$C_L = \int C_{\ell} dy \rightarrow \infty$$

Impossible

X



• Finite lift

• Velocity at tip is infinite

$$V \propto \Gamma \cdot \frac{1}{r}$$

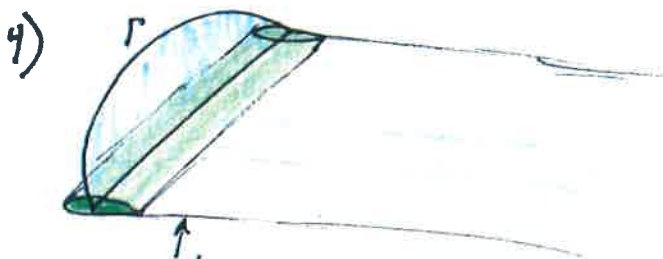
Reasonable mental model, but still wrong!



• Discrete vortices distributed along span

• Velocity at tip = 0 since $\Gamma_{tip} = 0$

• Trailing vortices induce α at wing which varies with span location.



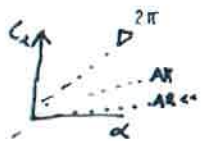
notice that the tip vortex has strength of zero!

• Continuous distribution of vorticity along span

✓

Analysis of Prandtl Lifting Line Theory:

The elliptical lift distribution is optimal for minimizing induced drag.



$$C_{L\alpha} = \frac{C_{L\alpha_{2D}}}{1 + \frac{C_{L\alpha_{2D}}}{\pi AR}}$$

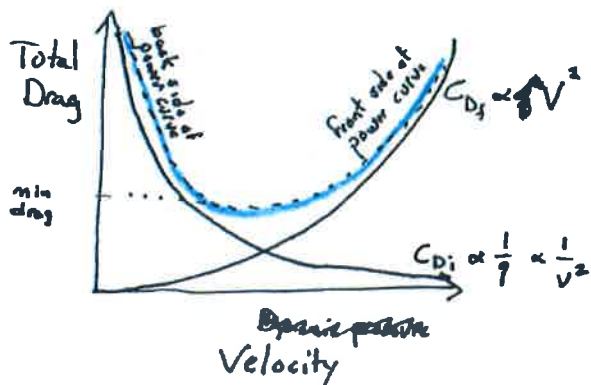
Reducing AR reduces lift slope to away from 2D value.

$$C_D = \frac{C_L^2}{\pi AR e}$$

where $e=1$
for elliptic
only

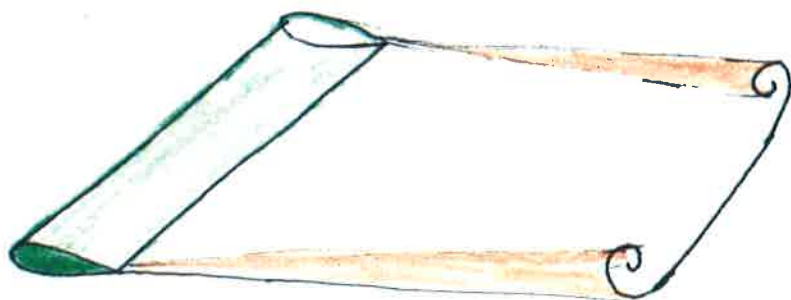
Induced drag depends on lift squared and the inverse of AR.

$$= \frac{\left(\frac{L}{b}\right)^2}{\rho g \pi e} \approx C \frac{(\text{span loading})^2}{(\text{dynamic pressure})^2}$$



Non-Elliptical distributions have an "e" value are essentially a ratio of actual to elliptical performance

$$e_{\text{non-elliptical}} < e_{\text{elliptical}} = 1$$



Wake rollup

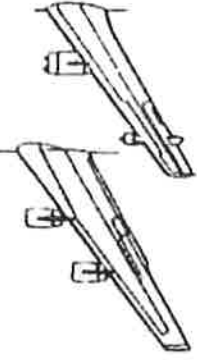
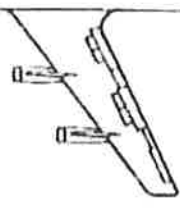




TYPE	B-47/B-52	367-80/KC-135	707-320/E-3A
FIRST FLIGHT	1947/1952	1954	1962
PLANFORM			
TYPICAL AIRFOIL			
$C_{L_{max}}$	1.8	1.78	2.2

Figure 26.4 - Trends in Boeing transport high-lift development.
Source: AGARD CP-365, paper no. 9


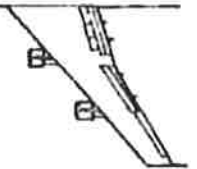
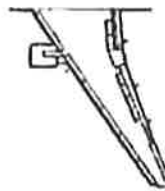



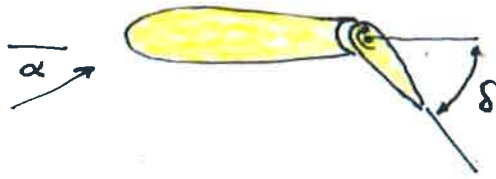
TYPE	727	747/E-4A	767
FIRST FLIGHT	1963	1969	1981
PLANFORM			
TYPICAL AIRFOIL			
$C_{L_{max}}$	2.78	2.45	2.45

Figure 26.5 - Trends in Boeing transport high-lift development - continued
Source: AGARD CP-365, paper no. 9

Hinge Moments

(Necessary for the upcoming flight control system portion of the class)



A moment is necessary to maintain the control surface at δ .

$$C_h = f(\alpha, \delta, Re, \text{gap, etc}) \text{ and possibly time}$$

$$H = C_h \cdot q \cdot S \cdot c \quad \left(\begin{array}{l} \leftarrow \text{chord aft of hinge} \\ \leftarrow \text{area aft of hinge} \end{array} \right) = C_h \cdot q \cdot c^2 \cdot W$$

The pilot is connected to the surface in some way



$$\delta = f(\text{stick angle})_{\delta_s}$$

$$F l_s \delta_s = H \delta_e \Rightarrow F = \left(\frac{\delta_e}{l_s \delta_s} \right) H e$$

How much force can a pilot exert? How long? $H \propto V^2$; pilot is limited human!
Constraints? Stick position, structural stresses, ...

Trim Tab



$$C_h = C_{h_0} + \underbrace{\frac{dC_h}{d\alpha}}_{\text{usually negative}} \alpha + \underbrace{\frac{dC_h}{d\delta_e}}_{\text{usually negative}} \delta_e + \underbrace{\frac{dC_h}{d\delta_t}}_{\text{usually positive}} \delta_t$$

For the pilot to have zero stick force, $C_h = 0$

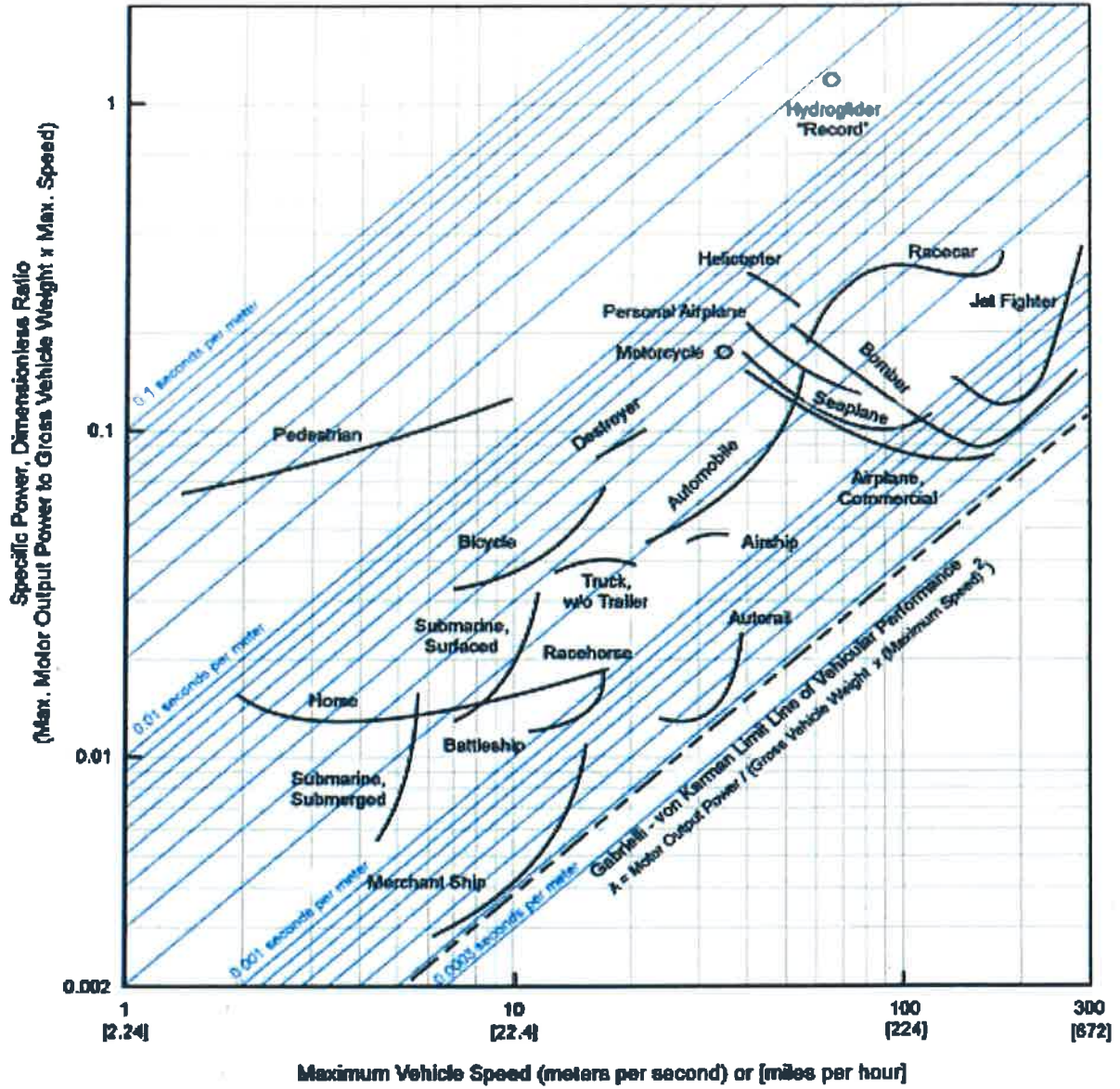
We (as pilots) can adjust δ_t to ensure $C_h = 0$ for a particular δ_e required at a particular α .

Stick free:

$$C_h = 0 = \left(C_{h_0} + \cancel{C_{h_{\delta_t}} \delta_t} \right) + \underbrace{C_{h_\alpha}}_{\text{usually negative}} \alpha + \underbrace{C_{h_{\delta_e}}}_{\text{usually negative}} \delta_e \Rightarrow \delta_e = - \frac{C_{h_\alpha}}{C_{h_{\delta_e}}} \alpha$$

The elevator "floats".
TEU $\propto \alpha$ increases

Von Karmen – Gabrielli Limit



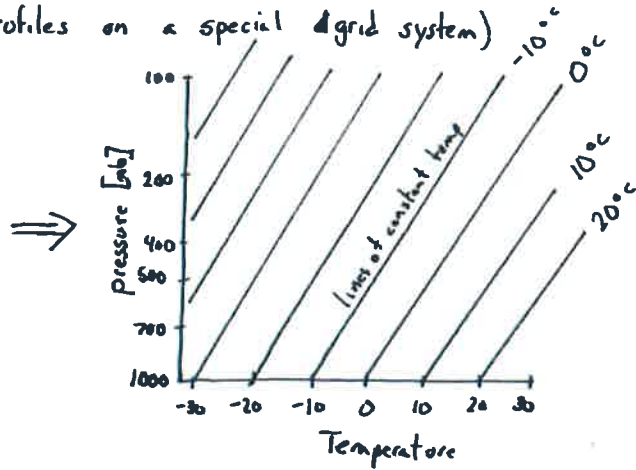
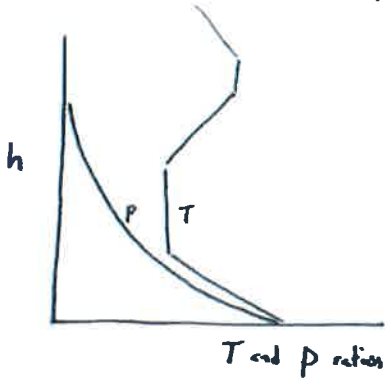
Lesson 4

Clouds

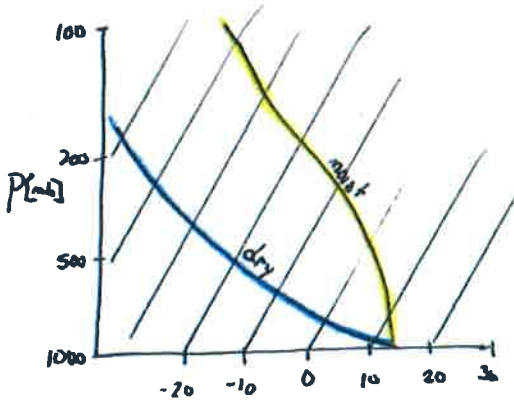
Why Clouds Form:

So far, we discussed only the dry-air lapse rate. Actual air has moisture.

skew T - Log P (draw pressure and Temp profiles on a special grid system)



Using a skew T-log P plot, lets look at the standard dry-lapse rate (dry)

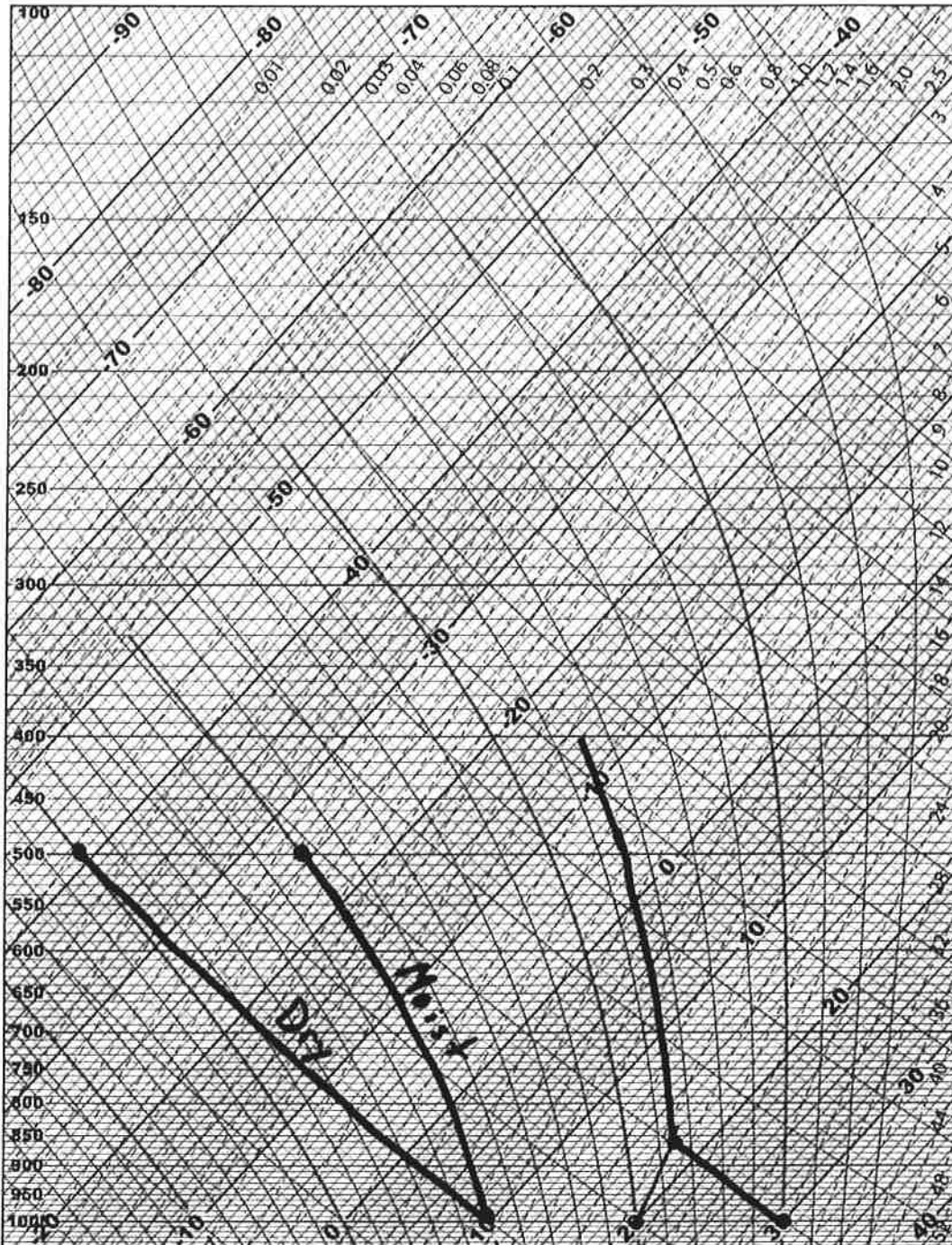


Changing the height to a log scale of pressure and tilting the temperature grid/coordinate allows us to see more of the atmosphere

Now, with water, the latent heat of water moving from ^{vapor} liquid to ^{liquid} vapor allows a slower cooling rate (Moist/Saturated lapse rate)

Once the moisture is condensed out (see previous Arden-Buck eqn) as the pressure drops, the moist lapse rate approaches the dry lapse rate.

Air ascends at the dry lapse rate until it becomes saturated (condenses), at which it ascends at the moist lapse rate



Skew-T / log p diagram for use with A First Course in Atmospheric Thermodynamics ©2008 by G.W. Petty

www.sundogpublishing.com

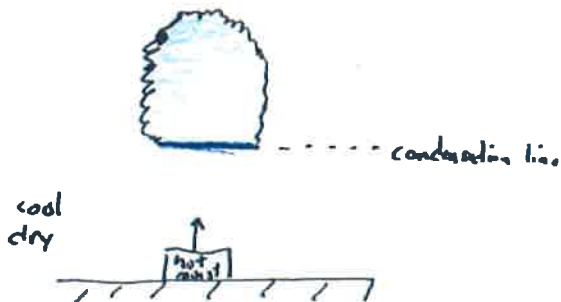
Example ① Lift moist vs dry air
 from $\underbrace{1000 \text{ mb}}_{\approx \text{SSL}}$ to $\underbrace{500 \text{ mb}}_{18000 \text{ ft}}$

Example ② Surface temp 30°C , Dew point 20°C
 Clouds form when condensation begins
 at intersection of dewpoint and dry lapse
 around $850 \text{ mb} \approx 4500 \text{ ft}$

Atmospheric Stability

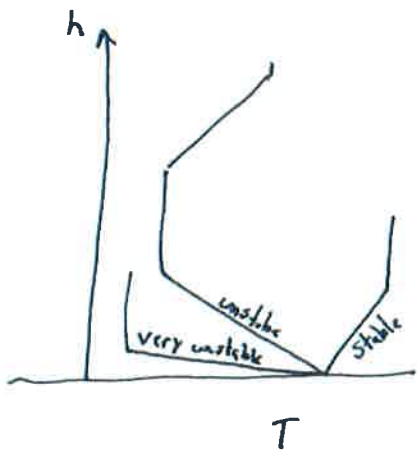
The driving force for cloud generation is density induced buoyancy. The hot moist air has a lower density and tends to rise within a cooler dry mass of air.

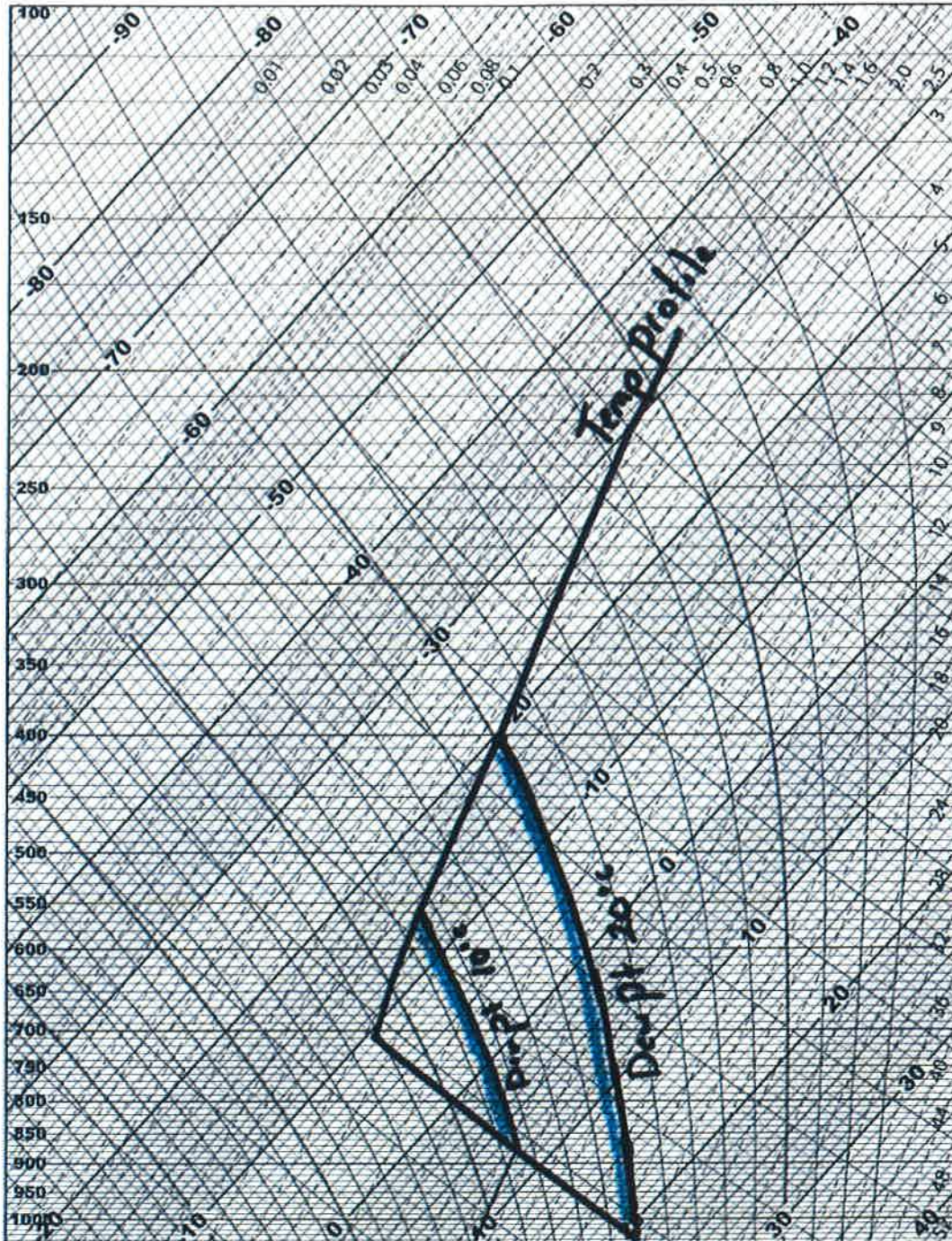
We saw that moist air rising has more energy (warmer) than dry air moved to the same height. This becomes an unstable system; As the moist air rises it tends to increase the driving force (density lowers).



For the standard atmosphere, where does the cloud stop rising? When the lapse rate changes to be more positive and eventually the density variation (and momentum) go to zero. The std atmosphere is naturally unstable to at least 36000 ft!

In reality, the convection/motion from the surface warms the upper atmosphere and reduces the lapse driving rate. Clouds ^{move} _{are a visible sign of} energy up!





Tuscaloosa

Op40 Analysis, valid 15-Jan-2016 04:00:00 (10.3nm/355° from TCL)

